

Preliminary Observations of Greenhouse Gases During the Summer and Winter ACT-America Campaigns

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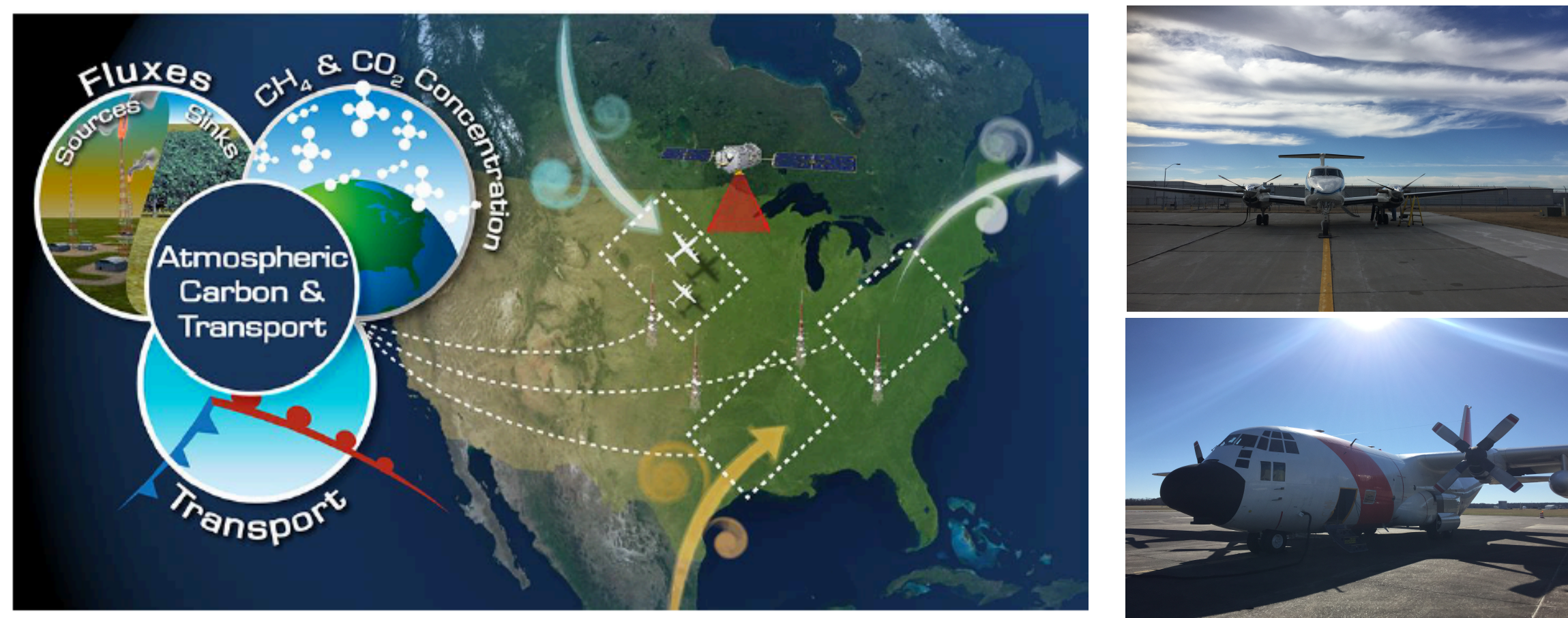
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1. Background: ACT-America goals

Carbon dioxide (CO₂) and methane (CH₄) fluxes and transport are poorly quantified at regional to continental scales. As a consequence, atmospheric inversion models have difficulty quantifying the carbon cycle and monitoring emissions and regulatory efforts at these scales. ACT-America is a multi-year, multi-campaign study that aims to provide high-density aircraft observations of regional and seasonal CO₂ and CH₄ across the eastern U.S. to a) reduce atmospheric transport model uncertainty, b) improve model prior estimates of carbon fluxes, and c) evaluate the spatial variability of satellite column CO₂ observations.



Figures 1-3. Figure 1 Left: Three ACT-America regions in the eastern U.S. used study carbon transport and fluxes using the NASA B-200 (Figure 2, top right) and NASA C-130 aircraft (Figure 3, bottom right). Flight days are separated into "fair weather", "frontal passage" and "O-CO₂ underpass" scenarios to evaluate carbon fluxes, transport, and satellite CO₂ retrievals, respectively.

2. NOAA flask measurements: sample drying

The NOAA Carbon Cycle and Greenhouse Gas (CCGG) aircraft network collects vertical profiles of air samples in portable flask packages (PFPs) throughout the U.S., and has been in operation since 1992. These packages are used during ACT-America and consist of twelve 0.7L borosilicate glass flasks pressurized to 275 kPa for a total sample volume of 2.2L. Flasks are analyzed for greenhouse gases including CO₂ and CH₄, carbon monoxide (CO), and ~50 other halocarbons, hydrocarbons and isotopic ratios useful for attributing CO₂ and CH₄ to different source types in each ACT region.

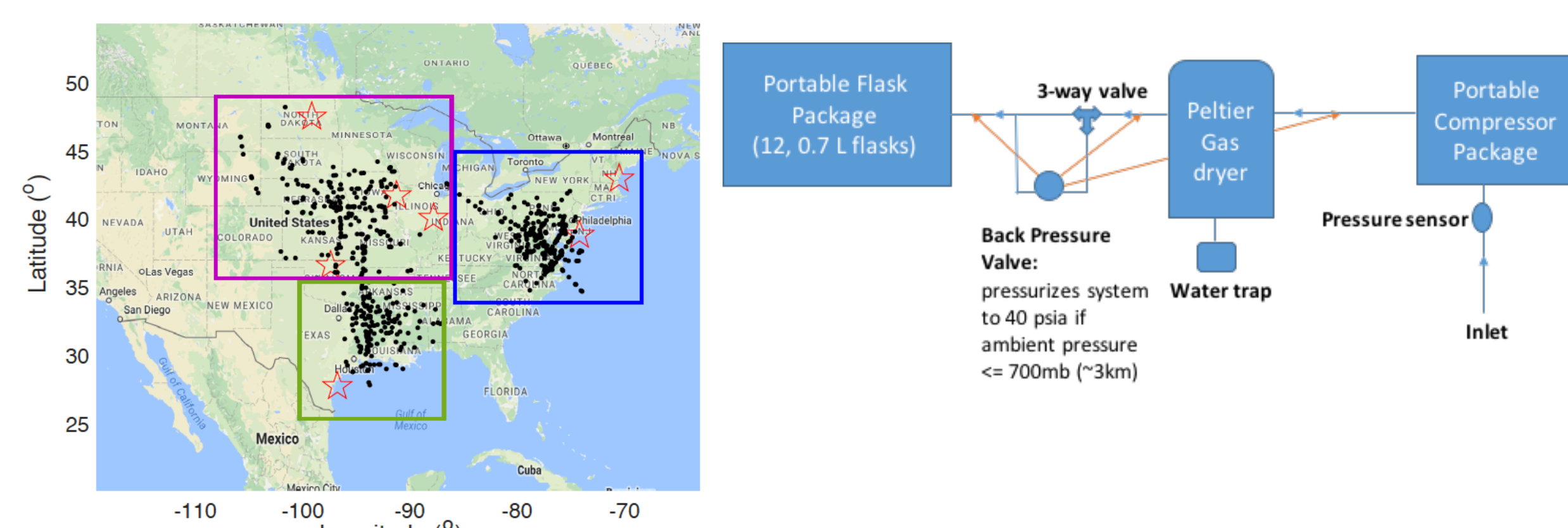


Figure 4 (left). Locations of all flask measurements made during summer 2016 and winter 2017. NOAA CCGG Network sites within the ACT domain are indicated by red stars. **Figure 5 (right).** Schematic of flask measurement system with gas chiller. Lines are pressurized below ~3km to facilitate condensing water out of air samples.

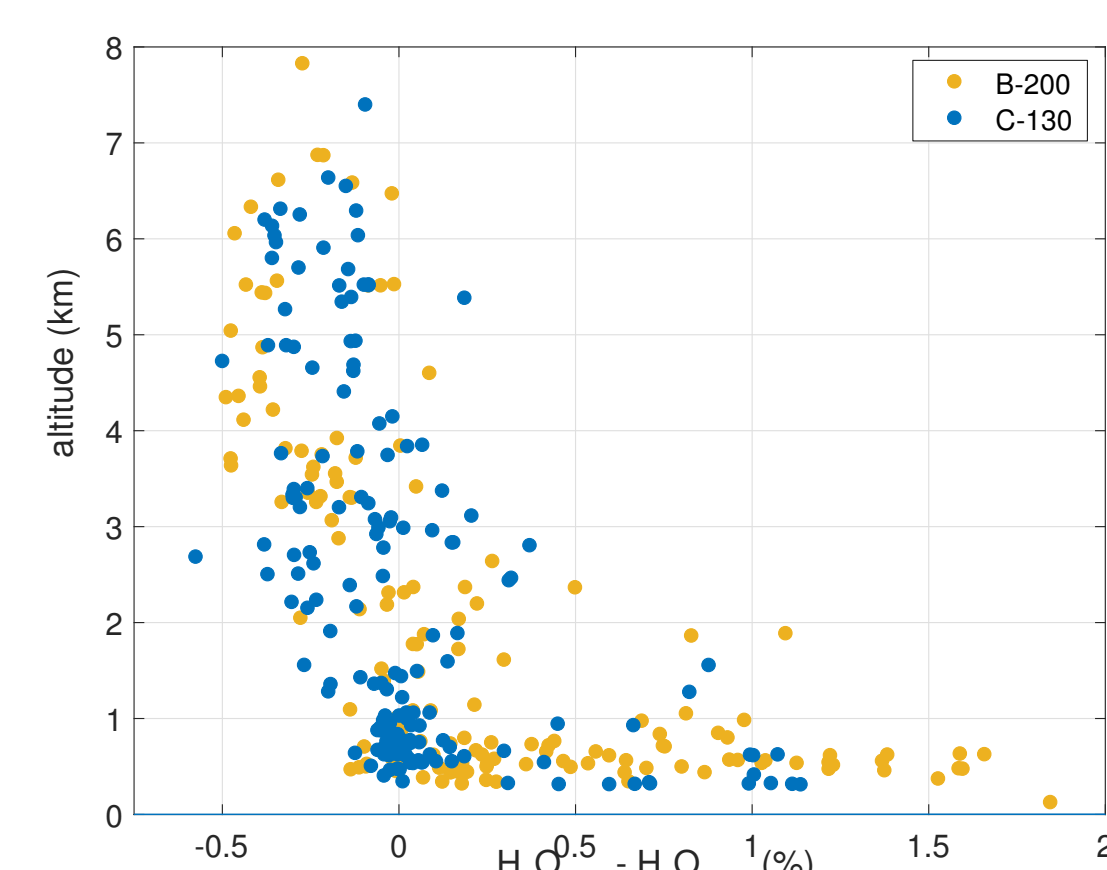


Figure 6. Difference between water vapor content (%) measured in ambient air and in air used to fill flasks post-drying on the B-200 and C-130 vs. altitude. Air samples are dried below ~3 km.

Flasks are pre-filled and flushed before pressurization to eliminate surface artifacts. Samples are also dried to ~0.5% water vapor using a Peltier gas chiller that is pressurized at altitudes < 3km to condense water vapor out of air samples. This drying helps to decrease the effects of water vapor on CO₂ measurements.

3. Flask vs. in situ data: improved agreement with flask drying

Comparisons between flask samples and continuous measurements made in situ using a Picarro cavity ring-down spectroscopy (CRDS) analyzer can help to identify uncertainties in greenhouse gas measurement methods. Generally, in situ-flask offsets were within the World Meteorological Organization (WMO) compatibility targets of 0.1 ppm for CO₂, 2 ppb for CH₄, and 2 ppb for CO. However, larger observed offsets may be due to uncertainties in flask fill times or systematic errors in CRDS measurements.

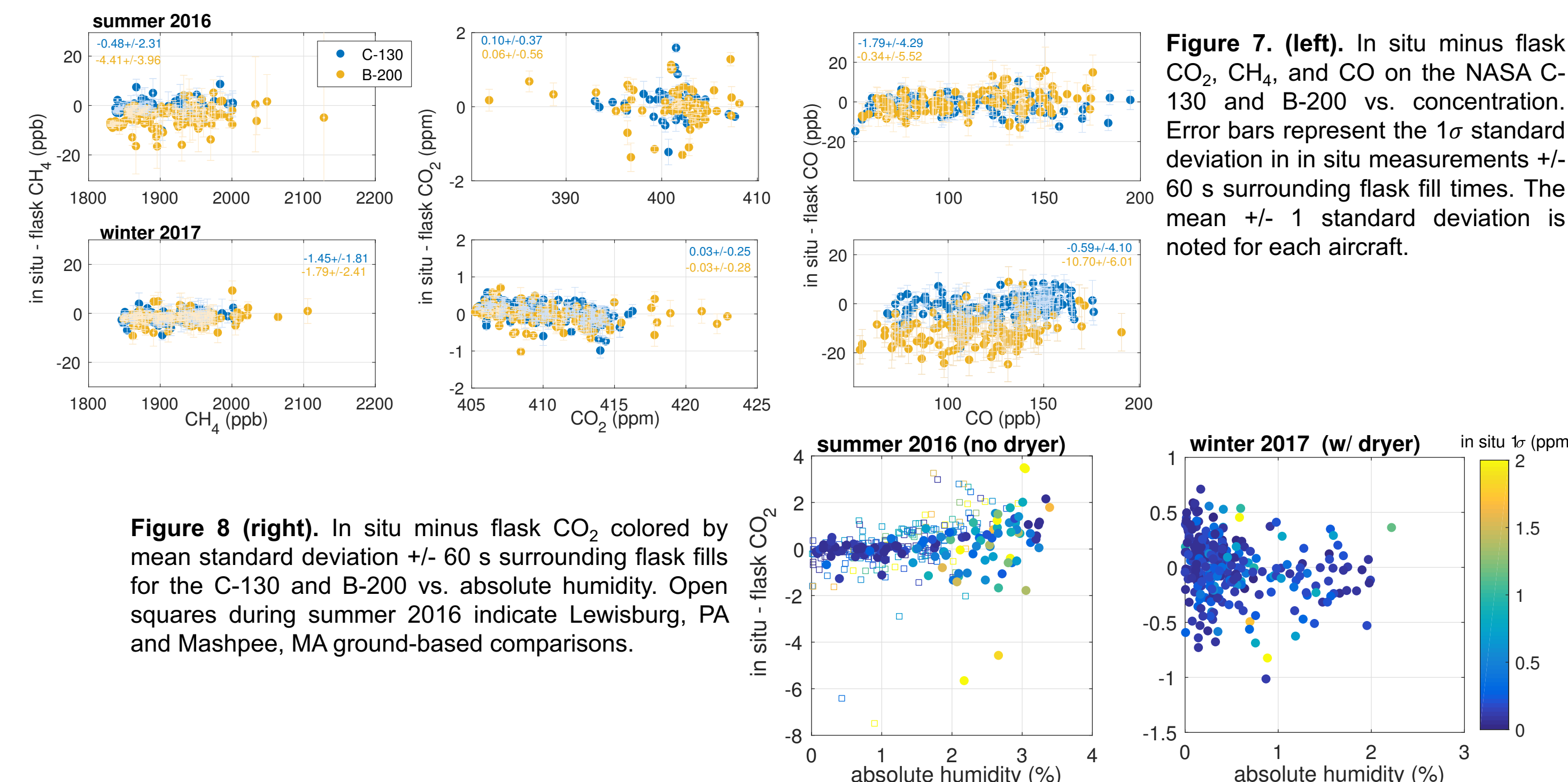


Figure 7 (left). In situ minus flask CO₂, CH₄, and CO on the NASA C-130 and B-200 vs. concentration. Error bars represent the 1 σ standard deviation in in situ measurements \pm 60 s surrounding flask fill times. The mean \pm 1 standard deviation is noted for each aircraft.

Figure 8 (right). In situ minus flask CO₂ colored by mean standard deviation \pm 60 s surrounding flask fills for the C-130 and B-200 vs. absolute humidity. Open squares during summer 2016 indicate Lewisburg, PA and Mashpee, MA ground-based comparisons.

Depletion in flask CO₂ and increased scatter relative to the CRDS data was observed above 1.5% absolute humidity during summer 2016. Similar results were observed for ground-based sites in Lewisburg, PA and Mashpee, MA when flask drying was not enabled. Drying flask samples during winter 2017 helped to decrease this water vapor effect on flask CO₂ measurements.

4. Greenhouse gas gradients across fronts: Higher warm sector CO₂, lower warm sector CH₄ during summer

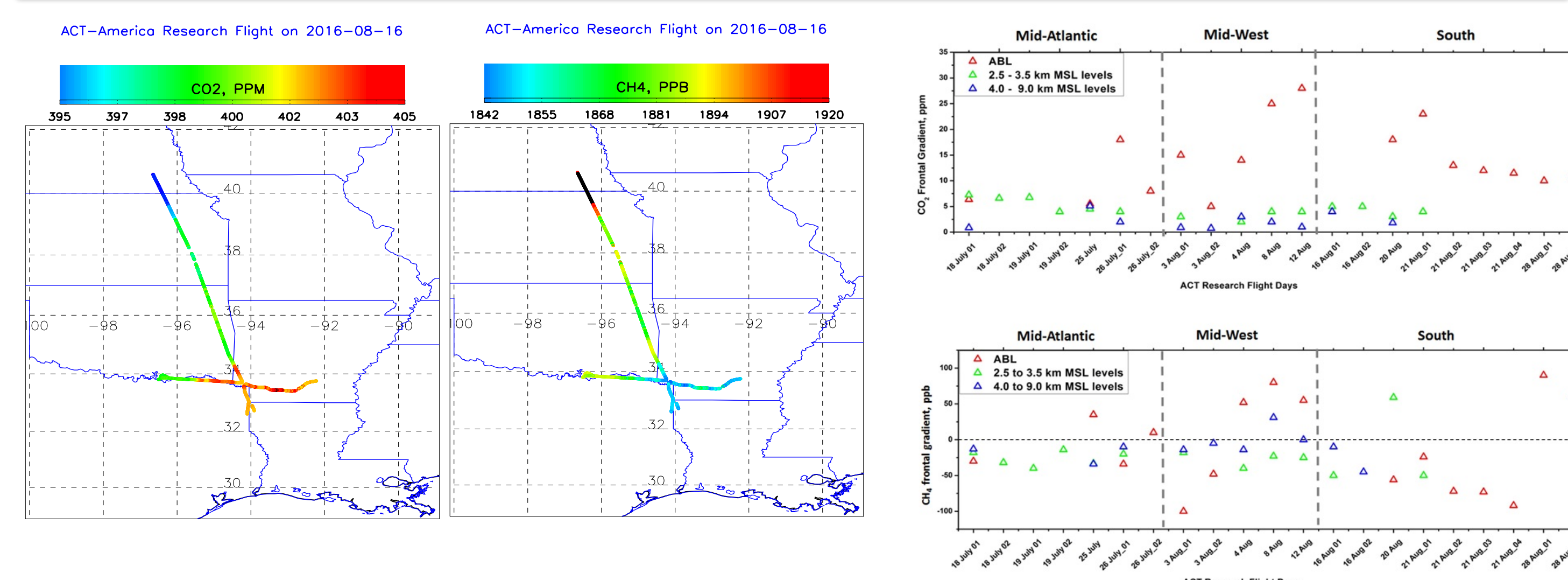
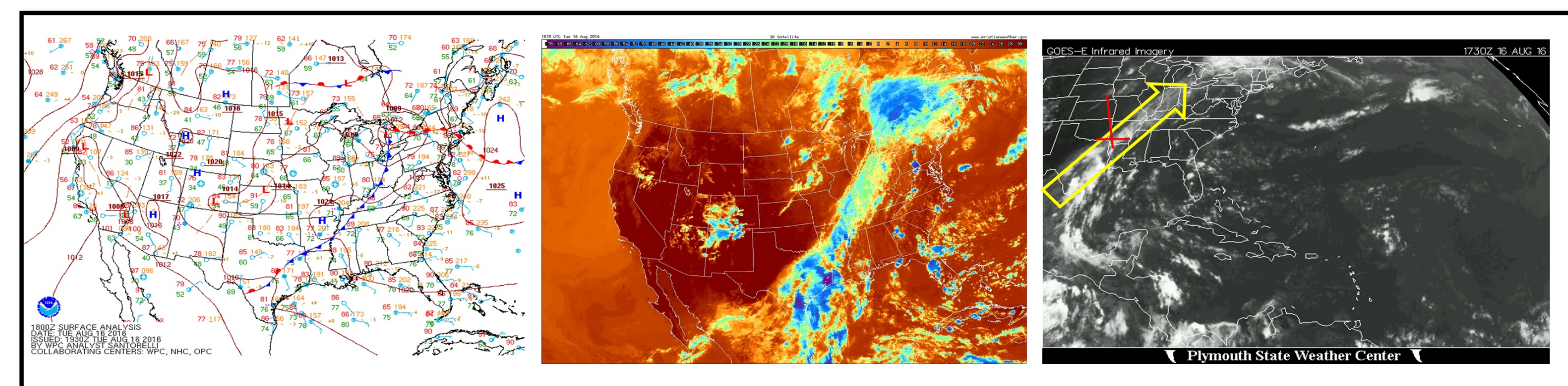


Figure 9 Top 3-panel: A cold front approached Shreveport, LA on Aug. 16, 2016. Moist air indicated by infrared and water vapor imagery propagated in a southwesterly fashion ahead of the front. **Figure 10 Bottom 3-panel.** Continuous in situ CRDS CO₂ and CH₄ gradients observed across the cold front on Aug. 16 2016. Regional gradients for all summer 2016 frontal passage flights in CO₂ and CH₄ colored by altitude (far right).

Continuous measurements of greenhouse gases across a cold front located near Shreveport, LA on August 16, 2016 indicated higher CO₂ concentrations in the warm sector than in the cold sector of the front. Conversely, observed CH₄ was lower in the warm sector than in the cold sector on this day. This general pattern was observed for both CO₂ and CH₄ across fronts during summer 2016 with larger variability observed in the atmospheric boundary layer (ABL) than in the free troposphere (FT).

5. CO₂ and CH₄ seasonal variability: comparisons to Mauna Loa, HI

Seasonal CO₂ variability depends on biogenic sources and sinks (summertime photosynthesis and wintertime respiration) and anthropogenic sources while seasonal CH₄ variability is driven by anthropogenic and natural sources and chemical oxidation. Observed greenhouse gas deviations from trends at Mauna Loa, HI were largest in the atmospheric boundary layer, with smaller deviations observed in the free troposphere.

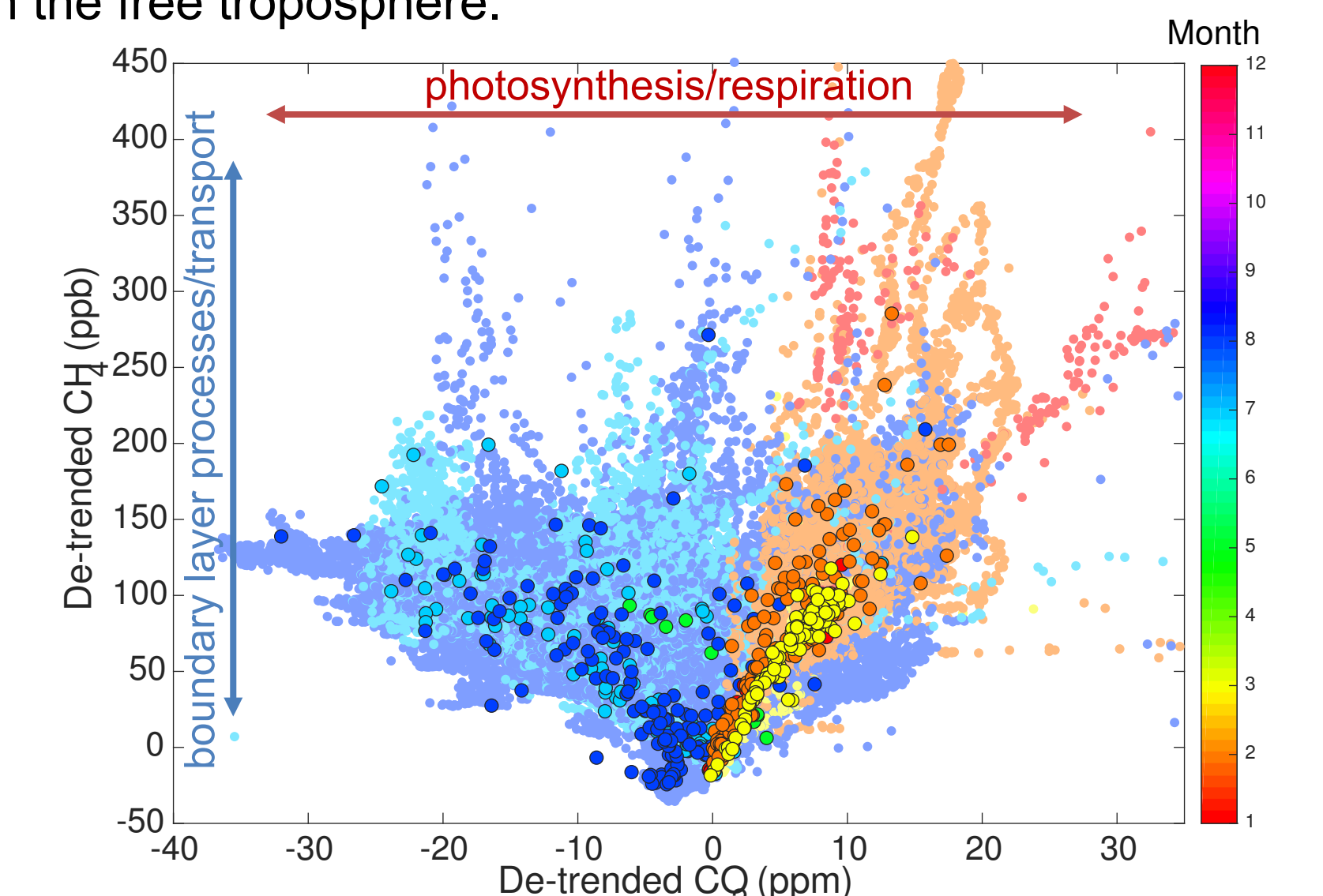


Figure 11. Flask CH₄ and CO₂ (bold circles) minus the observed CH₄ and CO₂ trend in Mauna Loa, HI for 2016-2017 colored by month. Faded background points indicate de-trended in situ CH₄ vs. CO₂ measured during the ACT-America campaign.

6. Correlations with anthropogenic species: strong wintertime correlations with urban tracers

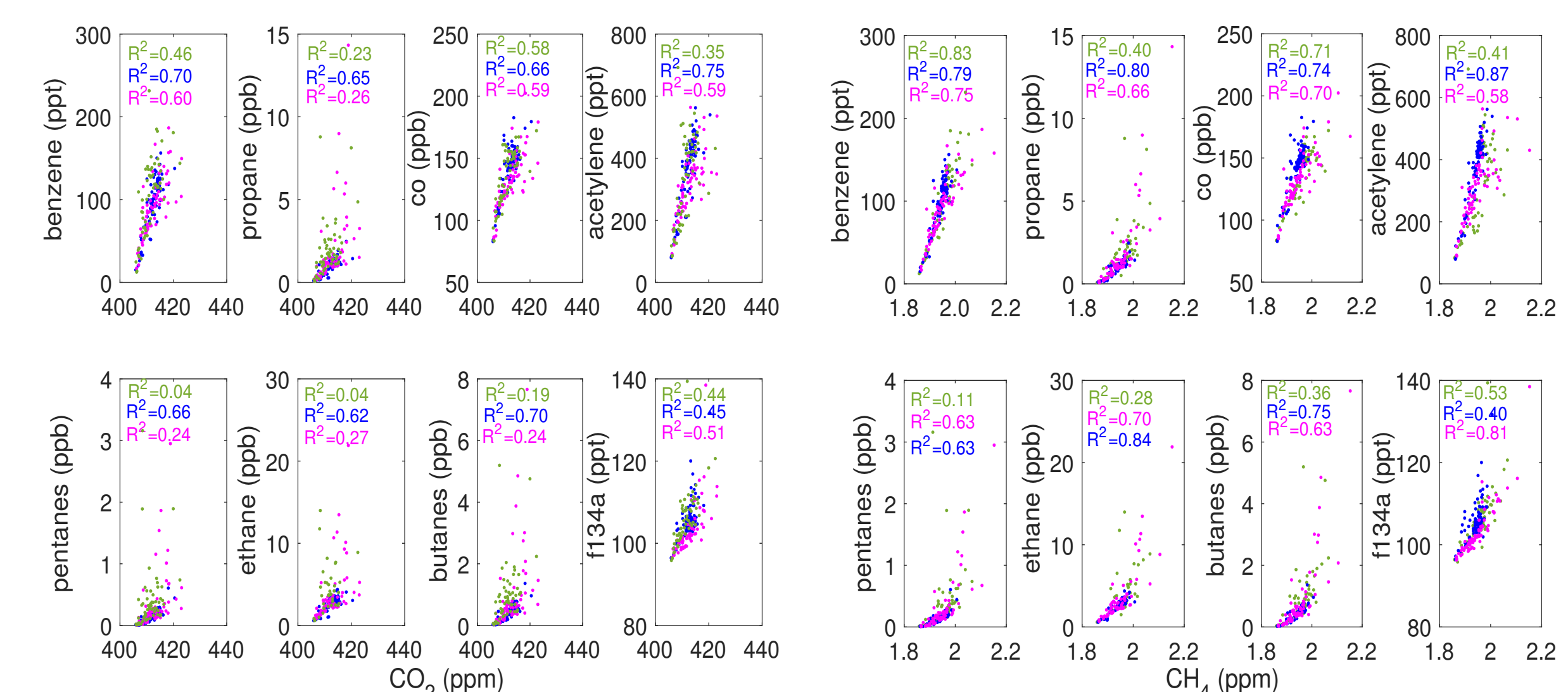


Figure 12. Lower-tropospheric (< 3 km agl) CO₂ and CH₄ correlations with various anthropogenic trace species in each ACT region during winter 2017. Colors indicate regions outlined in Fig. 4, with R² values indicated for each region.

Flask hydrocarbon species can be used to separate natural from anthropogenic sources of CO₂ and CH₄ and distinguish source signatures of atmospheric carbon. Preliminary observations indicate strong wintertime correlations between CO₂, CH₄ and urban tracers such as CO, acetylene (vehicular) and benzene (vehicular) throughout the ACT domain, with some of the highest R² values calculated in the northeastern United States.

7. Summary and future work

- Drying of flask samples helps to decrease observed water vapor - CO₂ artifacts
- ACT-America enables study of how frontal systems modify greenhouse gases
 - Investigate the causes for observed CO₂, CH₄ gradients across fronts
- Flask measurements help to identify sources and sinks of carbon, separate biogenic and anthropogenic contributions to total CO₂
 - Analyze back trajectories to evaluate air mass origin and source signatures of greenhouse gases
 - Analyze high density ¹⁴C₂ flask data to derive fossil fuel and biogenic contributions to total measured CO₂



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